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# ASSESSMENT OF IMPACTS OF CLIMATE CHANGE ON TROPICAL CYCLONE FREQUENCY AND INTENSITY IN THE TYPHOON COMMITTEE REGION

(Item 5 of the Provisional Agenda – Committee's Activities during 2009)

Submitted by Expert Team

**Action Proposed** 

The Committee is invited to:

- (a) Analyze and discuss the contents of the document;
- (b) Approve its publication after making the necessary changes

# Assessment of impacts of climate change on tropical cyclone frequency and intensity in the Typhoon Committee region

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# Foreword

(to be provided by the Typhoon Committee)

# Preface

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# Acknowledgements

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#### **Executive Summary**

There have been significant interdecadal and interannual fluctuations in the frequency of tropical cyclone (TC) formation and occurrence over the western North Pacific (WNP) in the last 50 years or so. Different interpretations on the trend of TC frequency over WNP have been reported, depending upon the best track dataset, analysis period, TC classification, etc. used. Thus, based on available publications, we cannot conclude whether there is a long term trend in the TC frequency over WNP. In view of this, an additional analysis utilizing 5 different best track datasets with data up to 2008 and allowing adjustments for the difference in averaging period between datasets has been conducted. The results suggest that most of the best track datasets depict either a decreasing trend or no trend in the annual number of TCs (tropical storm or above) and typhoons in WNP.

It should be noted that there exist significant differences between the available TC best track datasets for WNP. Such discrepancies could be due to the difference in the implementation of analysis techniques and the definition in the maximum wind as well as the limitation in in-situ observations.

For TC intensity, differences in best track data sets available for WNP do not allow for a convincing detection of a long term trend in TC intensity change in this basin when compared with variability from natural causes.

The trend of the number of landfalling tropical cyclone varies from one region to the other. There is no significant linear trend in the frequency of landfalling TCs in Japan and the Philippines. The trends of landfalling TCs in China and Thailand are decreasing. The trend of TC influencing Republic of Korea is increasing in recent years, but it is not conclusive.

In China, there is a decreasing trend in the maximum intensity of landfalling TCs in recent years but the mean intensity of landfalling TCs has no trend. The extreme wind induced by tropical cyclone affecting China has a decreasing trend and the total amount and intensity of TC precipitation has no significant trend.

Looking into the future, majority of the climate models projects a reduction in the number of TCs in the WNP in different future greenhouse gas scenarios. While there are fewer studies on the change of TC intensity, some of the model projections suggest an increase in the

number of intense TCs in the WNP in a warmer climate. Although climate models could provide us with projections for the future changes in TC activity, there exists a variety of uncertainties and limitations in the climate modeling and associated downscaling methods which may affect the skill and reliability of the projections, in particular in regional scale.

#### 1. Introduction

Climate change has become a hot topic of discussion in recent years. Besides global warming patterns, the possible change in tropical cyclone activity is also a matter of great concern to the public and decision makers.

Tropical cyclone is one of the most destructive weather systems on earth. The western North Pacific (WNP) is the most active tropical cyclone basin in the world with an annual average of about 30 tropical cyclones. To understand the possible changes in the tropical cyclone activity in the WNP under the climate change situation is a high priority issue in the Typhoon Committee region from both scientific and socio-economical viewpoints.

The Typhoon Committee is an intergovernmental body established in 1968 under the auspices of the United Nations Economic Commission for Asia and the Far East (UNECAFE) and the World Meteorological Organization (WMO). The Committee's purpose is to promote and coordinate planning and implementation measures required for minimizing the loss of life and material damage caused by typhoons. It is currently composed of 14 Members: Cambodia; China; Democratic People's Republic of Korea; Hong Kong, China; Japan; Lao People's Democratic Republic; Macao, China; Malaysia; the Philippines; Republic of Korea; Singapore; Thailand; Socialist Republic of Viet Nam and the United States of America.

In the Forty-first Session of the ESCAP/WMO Typhoon Committee held in Chiang Mai, Thailand (19 to 24 January 2009), it was decided that an Expert Team to be formed to assess the change in frequency and intensity of tropical cyclones in the region and make suggestions/ interpretation on this matter for the policymakers and the public. This decision was based on the information provided by Members, the outcomes of the Working Group on Meteorology parallel session, and the "Integrated Workshop on Coping with Climate Change in the Typhoon Committee Area" held in Beijing, China (22 - 26 September 2008),.

An Expert Team which consists of five experts from China; Japan; Republic of Korea; United States of America; and Hong Kong, China as well as a coordinator from Macao, China was formed in mid-2009 under the Working Group of Meteorology. The Expert Team conducted a series of literature review and assessment work in 2009 and met during the Expert Team Meeting held in Macao, China (14-15 December 2009) to discuss the findings and assessments.

In this report, the assessments on the observed changes in the tropical cyclone frequency, intensity, and landfalling are discussed in Sections 2, 3 and 4 respectively. Results of climate model projections for the future tropical cyclone activity given by various research groups and the uncertainty associated with the assessment are respectively reported in Sections 5 and 6. The recommendations for future work on the subject are included in Section 7. Moreover, for readers' easy reference, Annex I provides a comparison of the tropical cyclone classification and Annex II lists the acronyms used in the report.

It should be noted that the review and assessment by the Expert Team were conducted within a limited time period. While the Expert Team Members have strived to conduct the assessment based on the best available publications/information, the literature review process does not mean to be completely exhaustive. There is the possibility that certain findings and publications have not been covered in this assessment.

#### 2 Tropical cyclone frequency

There exist significant interannual and interdecadal fluctuations in the tropical cyclone (TC) frequency in the western North Pacific (WNP) over the past approximately 50 years (Yumoto and Matsuura, 2001; Yeung et al., 2005, Chan et al., 2009). As far as trend analysis is concerned, previous studies have shown divergent results (including increasing, decreasing and no trend) depending on the best track dataset, analysis period, and TC classification. Thus, based on available publications, it is difficult to conclude whether there is a long term trend in the TC frequency over WNP. In view of this, an additional independent analysis utilizing latest available data up to 2008 from five best track datasets and allowing adjustments for the difference in averaging period between datasets has been conducted to assess the possible trend of TC frequency in this basin. Trend analysis results reported by various researchers for different TC classifications are summarized in the Sections 2.1 to 2.3. It should be noted that the maximum sustained wind in the RSMC-Tokyo and HKO datasets have been averaged over a 10-minute period, while it is averaged over a 2-minute period in the CMA dataset and over a 1-minute period in the JTWC dataset. TC classifications and wind averaging periods adopted by different centers are tabulated in Annex I. The results of the independent analysis using the latest best track datasets from 1945 to 2008 are discussed in Section 2.4.

#### 2.1 Intense typhoon

Using the tropical cyclone best track data of Joint Typhoon Warning Center (JTWC), Webster *et al.* (2005) found that between the two consecutive 15-year periods of 1975-1989 and 1990-2004, the percentage of intense typhoons (maximum sustained wind (1-minute average) > 114 kts, i.e. equivalent to Cat. 4 to Cat. 5 hurricanes in the classification of the Atlantic basin) in the WNP has increased from 25% to 41%. However, Wu *et al.* (2006) noted that the same analysis using the best track datasets of the Hong Kong Observatory (HKO) and Regional Specialized Meteorological Center Tokyo (RSMC-Tokyo) showed a decrease in the proportion of intense typhoon between the two periods. Wu *et al.* (2006) also pointed out that there was a noticeable difference in the number of intense typhoons between the best track datasets of JTWC, HKO and RSMC-Tokyo during the period from 1977 to 2004. The study of Song *et al.* (2009) using the datasets of JTWC, RSMC-Tokyo and China Meteorological Administration (CMA) showed that, for TCs simultaneously recorded by these three datasets from 1977 to 2007, JTWC dataset showed an increasing trend for the

frequency of intense typhoons in WNP but the datasets of CMA and RSMC-Tokyo revealed a decreasing trend and no trend respectively.

# 2.2 Typhoon

Wu *et al.* (2006) and Yeung *et al.* (2005) noted that, between 1961 and 2004, there was no linear trend in the number of typhoons in WNP using the best track data of HKO. Similarly, no trend was found in the study conducted by Japan Meteorological Agency (JMA) using RSMC-Tokyo data between 1977 and 2008 (JMA, 2009). However, Ma and Chen (2009) reported a decreasing trend using a longer period (1949-2008) of data from CMA.

# 2.3 Other tropical cyclone categories

By analyzing HKO's best track data, Yeung *et al.* (2005) suggested that the annual number of TCs (all TCs, including tropical depressions) occurring over the WNP has been decreasing from 1961 to 2004. A similar trend for TCs (all TCs, excluding tropical depressions) has also been identified in the analyses carried out by Ma and Chen (2009) using the best track data of the CMA. Ma and Chen (2009) also analyzed the trends of each TC category using the CMA dataset and showed that typhoon and severe tropical storm have been decreasing, but tropical storm has been increasing slightly.

The study of Yuan *et al.* (2008) revealed that, based on the best track data of JTWC, the number of tropical storms in WNP has an increasing trend from 1945 to 2006, but the number of TCs in other categories did not exhibit any significant trend.

# 2.4 Independent assessment on TC frequency in WNP

Figure 2.1(a) shows the variation in the annual number of TCs with intensity reaching tropical storm or above from 1945 to 2008 based on the best track datasets of JTWC, RSMC-Tokyo, HKO, CMA and International Best Track Archive for Climate Stewardship (IBTrACS). To account for the differences in the definition of maximum wind averaging period between datasets (CMA uses 2-minute average, JTWC uses 1-minute average, other centers use 10-minute average), an adjustment was also performed to align the TC intensity of these two datasets to 10-minute average based on the maximum wind speed conversion factors for different averaging periods documented in the relevant WMO guidelines (Harper *et al.*, 2009).

Figure 2.1(b) also shows the TC (tropical storm or above) frequency variation for the 10-minute average adjusted datasets. Similar plots for the frequency of typhoons are given in Figures 2.2(a) and (b). T-test has also been conducted to check the statistical significance (at 5% level) of each trend. The linear trends of the annual number of TCs (tropical storm or above) and typhoons for the five datasets based on all available data (1945-2008) are tabulated in Table 2.1. A similar trend comparison based on the same data sets for the period from 1977 to 2008 is given in Table 2.2.

As shown in Table 2.1, for statistically significant trends, CMA dataset shows a decreasing trend for all four combinations, HKO dataset has a decreasing trend for "All TCs", and IBTrACS dataset shows a decreasing trend for "Typhoons". JTWC and RSMC-Tokyo datasets have no statistically significant trend (at 5% level) for all combinations.

For the trend analysis based on a common period from 1977 to 2008, all datasets and combinations have no statistically significant trend (see Table 2.2).

As such, based on the available data from 1945 to 2008 and after adjusting for the difference in the maximum wind averaging period between datasets, most of the best track datasets depict either a decreasing trend or no statistical significant trend in the annual number of TCs (tropical storm or above) and typhoons in WNP.



Intensity Defined as 10-min Sustained Wind



Figure 2.1 Variation in the annual number of TCs (tropical storm and above) in WNP for different best track datasets as computed based on (a) original wind speeds of various best track datasets, (b) the 10-minute average adjusted datasets. The shaded area between grey lines denotes the range of frequency differences among different datasets.





Intensity Defined as 10-min Sustained Wind



Figure 2.2 Same as Figure 2.1, but for annual number of typhoons in WNP.

Table 2.1 Trends of annual no. of TCs in WNP based on different datasets for all available data up to 2008. Data in bold indicates the trend is statistically significant at 5 % level.

	Data Period	Original inter	sity	10-minute averaging adjusted intensity		
Datasets		All TC	Typhoone	All TC	Typhoons	
		(tropical storm or above)	ryphoons	(tropical storm or above)		
СМА	1949-2008	-0.77/decade	-0.89/decade	-0.98/decade	-1.15/decade	
JTWC	1945-2008	+0.55/decade	-0.06/decade	+0.19/decade	-0.25/decade	
RSMC-Tokyo	1951-2008	-0.34/decade	-0.39/decade*	-0.34/decade	-0.39/decade*	
НКО	1961-2008	-1.35/decade	-0.59/decade	-1.35/decade	-0.59/decade	
IBTrACS	1945-2008	-0.67/decade	-0.62/decade	-0.67/decade	-0.62/decade	

\* period from 1977 to 2008 as maximum sustained wind data in RSMC-Tokyo dataset only available since 1977

	Original in	tensity	10-minute averaging adjusted intensity		
Datasets	All TC	Tunhoono	All TC	Typhoons	
	(tropical storm or above)	ryphoons	(tropical storm or above)		
СМА	CMA -1.18/decade -0.22/decade		-1.59/decade	-0.79/decade	
JTWC +0.08/decade		+0.26/decade	-0.69/decade	+0.26/decade	
RSMC-Tokyo	-1.00/decade	-0.39/decade	-1.00/decade	-0.39/decade	
НКО	-0.63/decade	-0.22/decade	-0.63/decade	-0.22/decade	
IBTrACS -1.22/decade		-0.15/decade	-1.22/decade	-0.15/decade	

Table 2.2 Trends of annual no. of TCs in WNP based on different datasets from 1977 to 2008. No trend is statistically significant at 5 % level.

#### 3. Tropical cyclone intensity

By using the homogeneous record of the University of Wisconsin-Madison/National Climatic Data Center (UW/NCDC), Kossin *et al.* (2006) indicated that there was no significant trend in the power dissipation index (PDI) in the WNP from 1983 to 2004. The study by Wu *et al.* (2006) also found no trend in PDI or intense typhoons in WNP from 1965 to 2004 using best track datasets of HKO and RSMC-Tokyo. However, a similar analysis conducted by Wu *et al.* (2008) using the best track data from JTWC indicated an increasing trend in PDI over WNP from 1975 to 2004. Such a discrepancy due to the choice of dataset and analysis period reduces our confidence in the assessment of TC intensity trend.

Similar dependency on dataset was also observed in the analysis of TC days. Kamahori *et al.* (2006) found that, based on RSMC-Tokyo best track dataset, there was a substantial decrease in the number of TC days for intense typhoons over the WNP between the periods 1977-1990 and 1991-2004. However, this result differed from the one using the JTWC dataset which showed an increasing trend in the number of TC days for intense typhoons. Such a disagreement among dataset in WNP could be attributed to different implementation of the Dvorak technique based on satellite based measurements and maximum wind definition (Kwon *et al.*, 2006; Yu *et al.*, 2007).

The study of Chan (2009) indicated that, in WNP, the thermodynamic factors such as maximum potential intensity (MPI) do not seem to have any appreciable contribution towards the variability of the annual occurrence of intense typhoons. Dynamic factors such as vertical wind shear may be the determinant in this basin. This implies that although global warming could likely increase the thermodynamic energy available in the atmosphere, such an increase does not necessarily imply a concomitant increase in the number of intense TCs. Until we can demonstrate that the dynamic factors will also become more favourable for TC intensification, it remains uncertain whether the frequency of occurrence of intense TCs will increase under a global warming scenario.

The other metrics of intensity such as TC precipitation still require further research, even though there is general support on the increasing tendency of moisture and associated precipitation in a warming climate on a global scale scenario (Trenberth, 1999 and 2008).

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#### 4. Landfalling

#### 4.1 Landfalling TC frequency

Chan *et al.* (2009) conducted a study on TC (TS and above) landfalling trends of 3 different sub-regions in East Asia, namely South (south China, Vietnam and the Philippines), Middle (east China), and North (Korean Peninsula and Japan) by using the JTWC best track data. Their results showed that none of the time series of annual number of landfalling TCs shows a significant linear trend. They also suggested that global warming has not led to a higher frequency of landfalling TCs or typhoons in any of these regions in Asia.

The trends for landfalling TCs in different countries/places vary from one to the other. The frequency of TCs that hit Thailand had a decreasing trend since mid-1960s (Thai Meteorological Department, 2009). In the Philippines, the number of landfalling TCs possessed an increasing trend in Visayas, a decreasing trend in Mindanao, and no significant trend in Luzon (PAGASA, 2008). There was no significant trend in the number of landfalling TCs in Japan (JMA, 2009).

Park *et al.* (2006) showed that the frequency of typhoons influencing Korea peninsula was increasing in recent years, but not conclusive due to limited time record of 50 years (1954-2003).

For China, several studies have been conducted to investigate the trends of landfalling TCs. Their results showed that the number of TCs landfalling in China was decreasing (Wang and Ren, 2008; Hu *et al.*, 2008; Yang *et al.*, 2009). The study of Yang *et al.* (2009) pointed out that the TC landfalling frequency in southern China was decreasing but the frequency trend of those landfalling over East China was not obvious. There was no trend in the total number of typhoons landfalling over all of China. Yang *et al.* (2009) also found that the locations of landfalling TCs over China were tending to approach the area of 23~35°N (Figure 4.1), which are consistent with Cao *et al.* (2006), who suggested that the locations of landfall were tending to approach the central part of the China's coastline.

Regarding the interannual and interdecadal variations in TC tracks, Kim *et al.* (2005) examined the variation in TC tracks in summertime (July-September) over WNP by analyzing the TC passage frequency with empirical orthogonal function (EOF) analysis method. They

identified three leading modes. Two tropical modes represented the long term trend and the relationship with ENSO and the other mode oscillated between south of Korea and southeast of Japan with an interannual time scales (so-called East Asia dipole pattern (EADP)). Liu and Chan (2008) also studied the TC occurrence pattern over WNP and found a significant interdecadal variation.

### 4.2 Landfalling TC intensity

The study of Yang *et al.* (2009) showed that the percentage of intense typhoons landfalling in China showed an increasing trend from 1949 to 2006. However, Wang and Ren (2008) suggested that maximum landfall intensities in China were stronger during 1950s~1970s but decreased in recent years. Wang an Ren (2008) also showed the mean landfall intensity didn't depict any linear trend during 1951~2004.

Ying *et al.* (2009a) and Ying and Chen (2009) examined the changes of tropical cyclone impact on China and different areas of China (south China, East China, Northeast China and China inland) using the CMA's best track and TC wind and precipitation observation datasets. They found that there was no trend in the frequency of TCs affecting China and these four sub-regions. On the regional scale, the extreme wind induced by TCs affecting China had an overall decreasing trend but the total amount and intensity of TC precipitation had no significant trend. However for individual stations, those with a decreasing trend of storm winds were mainly located near the southern part of China's coastline, and those stations with increasing trends in rainfall intensity were mainly located in the southeastern area of China.



Figure 4.1 The south most (red line) and north most (blue line) locations of TCs making landfall in China (in latitudes) (Yang *et al.*, 2009).

#### 5. Future projections

Different research groups have conducted model simulations using climate models, driven by different future greenhouse gas and sea surface temperature rise scenarios, to project the future change in tropical cyclone activities in the 21<sup>st</sup> century.

#### 5.1 Frequency

The projections given by the studies of Sugi *et al.* (2002), McDonald *et al.* (2005), Gualdi *et al* (2008) and Sugi *et al.* (2009) using the atmospheric general circulation models revealed that there would be a significant reduction in the frequency of TCs in the WNP basin under the global warming situation. Simulations using high resolution climate model (Oouchi *et al*, 2006; Bengtsson *et al.*, 2007) also suggested a significant decrease in the number of tropical cyclones in WNP. Moreover, Yokoi and Takayabu (2009) have examined the global warming impacts on TC genesis frequency over WNP projected by five CMIP-3 atmosphere-ocean coupled general circulation models which could realistically simulate horizontal distribution in tropical cyclogenesis over the WNP. All five models projected increasing trends of frequency in the central North Pacific and decreasing trends in the western part, with a maximum drop over the South China Sea (see Figure 5.1).



Figure 5.1 Multi-model ensemble of projected global warming impact on the genesis frequency. Yellow(blue) tone: all 5 models project increase (decrease) trends. (extracted from Yokoi and Takayabu, 2009)

However, Caron and Jones (2008) indicated that annual frequency in WNP would increase. The study of Stowasser *et al.* (2007) also projected the total number of cyclones in WNP would be slightly higher in the global warming climate than at present, but the overall increase is not statistically significant. This contrasts with the results from the above mentioned high resolution global model projections which found substantial reduction in the number of tropical cyclones in the global warming climate.

Zhao *et al.* (2007) has conducted a review on the projections of TC activity over WNP in the 21<sup>st</sup> century provided by 11 different global climate models under different greenhouse gas emission scenarios. They indicated that most of the models projected that the number of TCs in the WNP may decrease in the 21<sup>st</sup> century.

Summary of the projections of TC frequency change is given in Table 5.1.

#### 5.2 Intensity

For TC intensity projections, fewer studies have been conducted for WNP when compared with TC frequency projections. Stowasser *et al.* (2007) found that, in the 21<sup>st</sup> century, the average intensity and the number of intense storms over the entire WNP region rise significantly in the global warming simulation. The study of Knutson and Tuleya (2004) also projected the mean TC intensity would increase in WNP. Bengtsson *et al.* (2007) found an increase in the number of intense TCs for the Northern Hemisphere. Oouchi *et al.* (2006) depicted that TC intensity in WNP would increase slightly in the future greenhouse-warmed climate, but the increase is not statistically significant. The review by Zhao *et al.* (2007) suggested that the TC intensity may increase in the 21<sup>st</sup> century.

However, the result from Sugi *et al.* (2002) indicated no significant change in the intensity in WNP in the 21<sup>st</sup> century. Hasegawa and Emori (2005) found that TC intensity would decrease on average but the tropical cyclone precipitation would increase due to the increase in atmospheric moisture content.

Summary of the projections of TC intensity change is given in Table 5.2.

Study Reference	Model Type / details	GHG	Global	WNP
Sugi at al., 2002 T106 AGCM		2xCO2	Decrease (-34%)	Decrease (- 66%) 105°E to 135°W
McDonald et al., 2005 N144 HadAM3		IS95a	Decrease (-6%)	Decrease (-30%)
Hasegawa and Emori, 2005	CCSR/NIES/FRCGC	2xCO2		Decrease (-4%)
	AGCM, T106			
Yoshimura <i>et al.,</i> 2006	T106 AGCM	2xCO2	Decrease (-9 to -18%)	
Oouchi <i>et al.,</i> 2006	TL959 (20km mesh)	A1B	Decrease (-30%)	Decrease (-40%) 0-45°N, 100-180°E
Stowasser et al., 2007	Downscaling	6xCO2		Increase slightly, but not statistically
	NCAR CCSM2			significant.
Bengtsson <i>et al.</i> , 2007	ECHAM5	A1B		Decrease (-20 to -28%)
	T213 (60km)			
	T319 (40km)			
Gualdi <i>et al.,</i> 2008	SINTEX-G (SXG)	2xCO2; 4xCO2	Decrease (-16 to -44%)	Decrease (-20%)
	AOGCM, T106			
Caron and Jones, 2008	CMIP-3 multi-model	A1B, A2, and B1		Increase (7 to 22%)
	Yearly Genesis			
	Parameter (YGP)			
Sugi <i>et al.</i> , 2009	TL959 (20km mesh)	A1B	7 experiments decrease	5 experiments decrease (-26 to -38%)
	TL319 (60km mesh)		(-20 to -30%)	2 experiments increase (28 to 64%)
	8 experiments			
Yokoi and Takayabu, 2009	CMIP-3 multi-model	A1B, A2, and B1		Increase in the central North Pacific
				(5 <sub>Γ</sub> –20 <sub>Γ</sub> N, 150 <sub>Γ</sub> E−180 <sub>Γ</sub> ; CNP)
				Decreasing in the western part, with a
				maximum decrease over the South China
				Sea (10 ⊢ –25 Γ N, 110 Γ –120 Γ E; SCS).

Table 5.1 Summary of results for some of the climate model projections of TC frequency in the 21<sup>st</sup> century

Study Reference	Model Type / details	GHG	Global	WNP
Sugi <i>at al</i> ., 2002	T106 AGCM	2xCO2	No significant change	
Knutson and Tuleya, 2004	CMIP2	+1%/yr CO2		Increase (5 to 14%)
Hasegawa and Emori, 2005	CCSR/NIES/FRCGC	2xCO2		Decreases on average, but precipitation
	AGCM, T106			increase
Oouchi <i>et al.</i> , 2006	TL959 (20km mesh)	A1B	Increase (11%)	Increase (4%)
Stowasser et al., 2007	Downscaling	6xCO2		Average intensity and the number of
	NCAR CCSM2			intense storms rises significantly
Bengtsson <i>et al.</i> , 2007	ECHAM5	A1B	Increase (the Northern	
	T213 (60km)		Hemisphere)	
	T319 (40km)			

# Table 5.2 Summary of results for some of the climate model projections of TC intensity in the 21<sup>st</sup> century

#### 6. Uncertainties

#### 6.1 Best track datasets

Assessments on climate change and TC activity always encounter data problems, such as the inhomogeneity of data, validity of statistics or indices, and inconsistency between different data or phenomena. As suggested by Landsea et al. (2005), the quality of global TC best track data may be not good enough for assessing the trends. According to the comparison conducted by Yu et al. (2007), there was a large discrepancy in the best track data set between different national meteorological services and warning centres in WNP. Yu et al. (2007) suggested that the intensity differences among CMA, RSMC Tokyo and JTWC data sets were significant when the three are compared at the same maximum sustained wind averaging period (10-minute average). For 1988-2003, the mean difference between CMA and RSMC Tokyo were 0.6 m/s, and the difference between CMA and JTWC was 1.7 m/s. However, the maximum difference in the strength of the same TCs was more than 30m/s, and the number of typhoons in CMA best track datasets was greater than that of RSMC-Tokyo and JTWC (Zhang and Ying, 2009). Moreover, Nakazawa and Hoshino (2009) examined the tropical cyclone datasets of the RSMC-Tokyo and JTWC from 1987 to 2006 and identified significant differences in Dvorak parameters (CI and T numbers) in 1992-1997 and 2002-2005 (see Figure 6.1).

Ying *et al.* (2009b) evaluated the differences of the annual cycle of TC activities as revealed by JTWC, RSMC-Tokyo, and CMA datasets. As shown in Figure 6.2, the differences in the power spectra of pentad storm frequency among the three best track datasets increased in decadal and longer time periods. That is to say, significantly different long term variations will occur when different datasets are used to construct indices such as PDI and ACE. Song *et al.* (2009) also compared the best track datasets of JTWC, CMA and RSMC-Toyko and indicted that, for TCs simultaneously recorded by all three datasets from 1977 to 2007, JTWC dataset had more "intense typhoons" than that of the other two datasets, leading to an increasing trend in the annual frequency of intense typhoons and PDI.

The difference in best track data could be due to :

 different implementations of the Dvorak technique, the basis for TC intensity estimation at major centres after the deactivation of aircraft reconnaissance after 1987 (Kamahori *et al.*, 2006; Nakazawa and Hoshino, 2009);

- difference in the maximum wind definition. JTWC uses one-minute sustained wind, CMA uses 2-minute sustained wind and most of the other centers use 10-minute mean wind;
- (iii) limited in-situ observations available for validating TC intensity; and
- (iv) data inhomogeneity resulting from technological advancements, such as observational techniques, changes in equipment, increase in station number, etc.

# 6.2 Model uncertainties

The advance in climate modeling in the last decade has provided us with a useful tool to project the future climate changes under various greenhouse gas and aerosol emission scenarios. However, there still exists a variety of uncertainties and limitations in the climate modeling and associated downscaling methods which may affect the skill and reliability of the projections of TC activities. Some of these uncertainties and limitations include:

- (i) Future greenhouse gas and aerosol emission scenarios;
- Physics of climate modeling, (e.g., including model parameterizations, model response to greenhouse gases and aerosols, projections of sea surface temperature in different regions, air-sea interaction simulation capability, etc.);
- (iii) Downscaling techniques, (e.g., boundary conditions for dynamic downscaling, assumptions for statistical downscaling approach, etc.); and
- (iv) Model resolution, model-scenario combination and criteria for selecting simulated TC.



Figure 6.1 Temporal changes in the frequencies of the difference in CI-number between JTWC and JMA for 1987-2006. The color circles indicate the occurrence number at each estimated Dvorak parameters and the numbers in the red circles show the percentage of the occurrence at the point (Nakazawa and Hoshino, 2009).



Figure 6.2 The power spectra for pentad storm frequencies that derived from the CMA, RSMC Tokyo (JMA) and JTWC datasets with original wind speeds (Ying *et al.*, 2009a).

# 7. Recommendations for Future Work

The following areas could be considered for future research on the subject in the Typhoon Committee region:

- Explore the feasibility of setting up or adopting a homogenous and unified TC best track dataset for WNP.
- Conduct more detailed studies on the observed TC activity in coordination with the Typhoon Committee's Training and Research Coordination Group, including the standardization of metrics, such as maximum sustained winds/minimum central pressure and station observed wind/pressure/precipitation, which have not been covered in past studies.
- Understand and validate the projections of future TC activities given by the global climate models or coupled general circulation models.
- Study the future TC impacts (e.g., wind, precipitation, storm surge, landfalling process, etc.) using high resolution global climate models and formulate relevant climate change adaptation and mitigation strategies.
- Use extreme statistical methods instead of the normal distribution approach to study the changes in TC activities.
- Promote sharing of knowledge and expertise with other Tropical Cyclone regional bodies and programmes.

#### References

Bengtsson, L., K. I. Hodges, M. Esch, N. Keenlyside, L. Kornblueh, J.-J. Luo, & T. Yamagata, 2007 : How may tropical cyclones change in a warmer climate. Tellus, 59A, pp. 539-561

Cao, C., J. -Y. Peng, and J.-H. Yu, 2006: An analysis on the characteristics of landfalling Typhoon in China under Global Climate Warming. J. Nanjing Institute of Meteorology. 29(4): 455-461. (In Chinese with English Abstract)

Caron, L.-P., and C. G. Jones, 2008: Analysing present, past and future tropical cyclone activity as inferred from an ensemble of coupled global climate models. Tellus. 60A: 80-96.

Chan J. and M. Xu, 2009: Inter-annual and inter-decadal variations of landfalling tropical cyclones in East Asia. Part I: time series analysis, International Journal of Climatology, Vol 29, pp.1285-1293

Chan J., 2009: Thermodynamic control on the climate of intense tropical cyclones. Proceedings of the Royal Society A, Vol. 465, pp 3011-3021.

Gualdi, S., E. Scoccimarro, and A. Navarra, 2008: Changes in Tropical Cyclone Activity due to Global Warming: Results from a High-Resolution Coupled General Circulation Model , , J. Climate, Volume 21, Issue 20 (October 2008) pp. 5204–5228.

Harper, B. A., J. D. Kepert, and J. D. Ginger, 2009: Guidelines for converting between various wind averaging periods in tropical cyclone conditions, WMO, 53pp.

Hasegawa, A., and S. Emori (2005), Tropical Cyclones and Associated Precipitation over the Western North Pacific: T106 Atmospheric GCM Simulation for Present-day and Doubled CO2 Climates. SOLA, 1: 145-148. doi: 10.2151/sola.2005-038.

Hu, Y., L. Song, A. Liu, and W. Pan, 2008: Analysis of the Climatic Features of Landfall Tropical Cyclones in China during the Past 58 Years. Acta Sci. Natural UM Univ. Sunyatseni, 47(5): 115-121. (in Chinese with English abstract)

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JMA, 2009 : Climate Change Monitoring Report 2008. http://ds.data.jma.go.jp/tcc/tcc/products/gwp/gwp.html

Kamahori, H., N. Yamazaki, N. Mannoji, and K. Takahashi, 2006: Variability in intense tropical cyclone days in the western North Pacific. SOLA, 2: 104-107, doi: 10.2151/sola.2006-027.

Kim, J.-H., C.-H. Ho, C.-H. Sui, and S.-K. Park, 2005 : Dipole structure of interannual variations in summertime tropical cyclone activity over East Asia. J. Climate, Vol 18, pp.5344-5356.

Kossin, J. P., K. R. Knapp, D. J. Vimont, R. J. Murnane, and B. A. Harper, 2007 : A globally consistent reanalysis of hurricane variability and trends, Geophys. Res. Lett., 34, L04815, doi:10.1029/2006GL028836.

Knutson, T. R., & Tuleya, R. E., 2004 : Impact of CO2-induced warming on simulated hurricane intensity and precipitation: sensitivity to the choice of climate model and convective parameterization. J. Climate, Vol 17, pp.3477-3495.

Kwon, H. J., S.H. Won and S K Park, 2006 : Climatological Differences between the Two Typhoon Centers Tropical Cyclone Information in the Western North Pacific, Journal of the Korean Meteorological Society, Vol. 42 (3), pp.183-192.

Landsea, C. W., B. A. Harper, K. Hoarau, and J. A. Knaff, 2005 : Can we detect trends in externe tropical cylones? Science. 313: pp.452-454.

Liu, K.S. and Johnny C. L. Chan, 2008 : Interdecadal Variability of Western North Pacific Tropical Cyclone Tracks, J. Climate, Volume 21, Issue 17 (September 2008) pp. 4464–4476. http://ams.allenpress.com/perlserv/?request=get-abstract&doi=10.1175%2F2008JCLI2207.1

Ma, L.P. and L.S. Chen, 2009 : The relationship between global warming and the variation in tropical cyclone frequency over the western North Pacific. J. Tropical Meteorology, Vol 15, No. 1, pp.38-44.

McDonald, R.E., Bleaken, D.G., Cresswell, D.R., Pope, V.D., & Senior, C.A., 2005 : Tropical storms: representation and diagnosis in climate models and the impacts of climate change.

Clim. Dyn., 25: 19-36, DOI: 10.1007/s00382-004-0491-0.

Nakazawa, T. and S. Hoshino, 2009 : Intercomparison of Dvorak parameters in the tropical cyclone datasets over the western North Pacific. SOLA, 5, 33-36. doi:10.2151/sola.2009 009. http://www.jstage.jst.go.jp/article/sola/5/0/33/ pdf

Oouchi K., J. Yoshimura, H. Yoshimura, R. Mizuta, S. Kusunoki and A. Noda, 2006 : Tropical Cyclone Climatology in a Global-Warming Climate as Simulated in a 20 km-Mesh Global Atmospheric Model : Frequency and Wind Intensity Analysis, Journal of the Meteorological Society of Japan, Vol. 84, No. 2, pp. 259-276.

PAGASA, 2008 : Philippine tropical cyclone trends. "Integrated Workshop on Coping with Climate Change in the Typhoon Committee Area", Beijing, China, 22 - 26 September 2008.

Park, J-K., B.-S. Kim, W.-S. Jung, E.-B. Kim, and D.-G. Lee, 2006: Change in statistical characteristics of typhoon affecting the Korea Peninsula. Atmosphere. 16: 1-17.

Stowasser M., Y.Q. Wang and K. Hamilton, 2007. Tropical Cyclone Changes in the Western North Pacific in a Global Warming Scenario J. Climate , Vol 20 (11). <u>http://ams.allenpress.com/perlserv/?request=get-abstract&doi=10.1175%2FJCLI4126.1</u>

Sugi M., A. Noda and N. Sato, 2002 : Influence of the Global Warming on Tropical Cyclone Climatology : An Experiment with the JMA Global Model, Journal of the Meteorological Society of Japan, Vol. 80, No. 2, pp. 249-272.

Sugi M., H. Murakami and J. Yoshimura, 2009 : A reduction in global tropical cyclone frequency due to global warming. SOLA, Vol. 5, 000-000, doi:10.2151/sola. 2009-000.

Song, J.J., Y. Wang and L.G. Wu, 2009 : Trend discrepancies in western North Pacific tropical cyclones. J. Geophysical Research (accepted for publication).

Thailand Meteorological Department, 2009 : Member Report. "Integrated Workshop : Building Sustainability and Resilience in High Risk Areas of the Typhoon Committee", 14-18 September 2009, Cebu, Philippines. Trenberth, K. E., 1999 : Conceptual framework for changes of extremes of the hydrological cycle with climate change. Climatic Change, 42, 327-339.

Trenberth, K. E., 2008 : The Impact of Climate Change and Variability on Heavy Precipitation, Floods, and Droughts, in The Encyclopedia of Hydrological Sciences, edited by M. G. Anderson, John Wiley & Sons, Ltd., Chichester, UK. doi: 10.1002/0470848944.hsa211.

Wang, X., and F. Ren, 2008: Variations in frequency and intensity of landfall tropical cyclones over China during 1951~2004. Marine Forecasts. 25(1): 65-73. (in Chinese with English Abstract)

Webster, P. J., G. J. Holland, J. A. Curry, and H.-R. Chang, 2005: Changes in tropical cyclone number, duration, and intensity in a warming environment. *Science*. 309: 1844-1846.

Wu, M.C., K.H. Yeung & W. L. Chang, 2006. Trends in Western North Pacific Tropical Cyclone Intensity, EOS transaction, AGU, Volume 87, Number 48, p537-538, 28 November 2006. <u>http://www.weather.gov.hk/publica/reprint/r683.pdf</u>

Wu, L. G., B. Wang and S. A. Brauna, 2008 : Implications of tropical cyclone power dissipation index. Int. J. Climatol., Vol. 28, pp.727–731.

Yang, Y., M. Ying, and B. Chen, 2009: Climatic changes of landfall tropical cyclones in China over the Past 58 years. (in press, in Chinese with English Abstract)

Yeung, K. H., M.C. Wu, W.L. Chang & Y.K. Leung, 2005. Long-term Change in Tropical Cyclone Activity in the Western North Pacific, Scientific Assembly of International Association of Meteorology and Atmospheric Science (IAMAS) 2005, Beijing, China, 2-11 August, 2005. http://www.weather.gov.hk/publica/reprint/r601.pdf

Ying, M., Yang. Y. and B. Chen, and W. Zhang, 2009a: Climatic variation of tropical cyclones affecting china during past 50 years. (submitted, in Chinese with English abstract)

Ying, M., E.J. Cha, H.J. Kwon, and S.H. Won, 2009b: Comparison of three western North Pacific tropical cyclone best track datasets in seasonal context. (submitted)

Ying, M. and B. Chen, 2009: Climatic Trend of the Tropical Cyclones' Influences on China's Mainland as Revealed by the Wind and Precipitation Observations. Reported in WMO 2nd International Workshop on Tropical Cyclone Landfall Processes (IWTCLP-II). 19-23 October 2009, Shanghai, China.

Yokoi S. and Y. N. Takayabu, 2009 : Multi-model Projection of Global Warming Impact on Tropical Cyclone Genesis Frequency over the Western North Pacific, Journal of the Meteorological Society of Japan, Vol. 87, No. 3, pp. 525-538.

Yoshimura, J. and M. Sugi, 2006 : Influence of Greenhouse Warming on Tropical Cyclone Frequency, J. Meteorological Society, Japan, 84: 405-428. 2006. http://www.jstage.jst.go.jp/article/jmsj/84/2/84\_405/\_article

Yu, H., C. Hu, and L. Jiang, 2007: Comparison of three tropical cyclone intensity datasets. Acta Meteorologica Sinica, Vol. 21, No. 1, pp.121-128.

Yuan, J., A. Lin, and C. Liu, 2008: Change characters of tropical cyclones with different intensities over the western North Pacific during the last 60 years. Acta Meteorologica Sinica, 66(2) 213-223. (in Chinese with English abstract)

Yumoto, M. and T. Matsuura, 2001. Interdecadal Variability of Tropical Cyclone Activity in the Western North Pacific, Journal of the Meteorological Society of Japan, vol. 79, No. 1, pp 23-35.

Zhang, X., and M. Ying, 2009: Analysis the differences of three tropical cyclone best track datasets. Marine Forecasts, Vol.26 No. 3, pp.60-70 (in Chinese).

Zhao, Z.C., Y. Luo, X.J. Gao and Y. Xu, 2007 : Projections of typhoon changes over the western North Pacific ocean for the 21<sup>st</sup> century. Adv. Clim, Change Res., Vol 3, No. 3, pp.158-161. (in Chinese with English abstract)

Maximur	n Sustained V	Wind Speed					
at the centre of the tropical		tropical	Hong Kong	Mainland China	Japan	US Pacific	US Atlantic
cyclones			(10-minute average)	(2-minute average)	(10-minute average)	(1-minute average)	(1-minute average)
kts	km/h	m/s					
< 34	< 63	<17.1	Tropical Depression (TD)	Tropical Depression	Tropical Depression	Tropical Depression	Tropical Depression
34 – 47	63 - 87	17.2-24.4	Tropical Storm (TS)	Tropical Storm	Tropical Storm	Tropical Storm	Tropical Storm
48 – 63	88 – 117	24.5-32.6	Severe Tropical Storm (STS)	Severe Tropical Storm	Severe Tropical Storm		
64 90	119 140	22.7.41.4	Typhoon (T)	Typhoon	Typhoon :	Typhoon	Hurricane categories
04 - 80	110 - 149	52.7-41.4			64 – 84 kts	64-129kts	1: 64 – 82 kts
			Severe Typhoon (ST)	Severe Typhoon			
81 – 99	150 – 184	41.5-50.9					2: 83 – 95 kts
					Very Strong Typhoon		
			Super Typhoon (SuperT)	Super Typhoon	85 – 104 kts		3: 96 – 113 kts
>=100	>=185	>=51			Violent Typhoon		
					>=105 kts		4: 114 – 135 kts
						Super Typhoon: >= 130 kts	
							5: >135 kts

# Annex I Comparison of the Tropical Cyclone Classification

Note : the conversion between kts to km/h and kts to m/s may vary slightly subject to rounding practices and conversion factor decimal places.

# Annex II Acronyms

Acronyms	Full Name			
ACE	Accumulated Cyclone Energy			
AGCM	Atmospheric General Circulation Model			
AOGCM	Atmosphere-Ocean Coupled General Circulation Model			
APEC	Asia-Pacific Economic Cooperation			
СМА	China Meteorological Administration			
CMIP	Coupled Model Intercomparison Project			
TC(s)	Tropical Cyclone(s)			
EADP	East Asia Dipole Pattern			
ESCAP	UN Economic and Social Commission for Asia and the Pacific			
GHG	Greenhouse gas			
нко	Hong Kong Observatory			
IBTrACS	International Best Track Archive for Climate Stewardship			
JMA	Japan Meteorological Agency			
JTWC	Joint Typhoon Warning Center			
MPI	Maximum Potential Intensity			
NOAA	National Oceanic and Atmospheric Administration			
PAGASA	Philippine Atmospheric, Geophysical & Astronomical Services			
	Administration			
PDI	Power Dissipation Index			
RSMC-Tokyo	Regional Specialized Meteorological Center Tokyo			
TRCG	Training and Research Coordination Group			
WMO	World Meteorological Organization			
WNP	Western North Pacific			